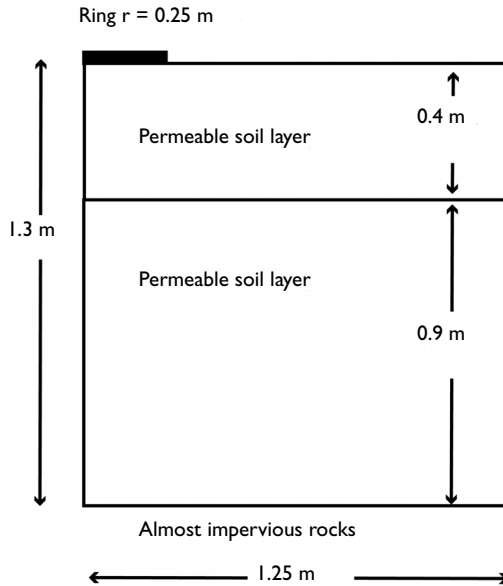


# Variably Saturated Flow and Transport—Sorbiting Solute

## Introduction

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In this example, water ponded in a ring on the ground moves into a relatively dry soil column and carries a chemical with it. As it moves through the variably saturated soil column, the chemical attaches to soil particles, slowing the solute transport relative to the water. Additionally the chemical concentrations decay from biodegradation in both the liquid and the solid phase. The inspiration for the problem comes from [Ref. 1](#).



*Figure 1: Geometry of the infiltration ring and soil column.*

The example uses the Richards' Equation interface to define nonlinear relationships with retention and permeability properties according to van Genuchten ([Ref. 2](#)). In the Richards' Equation interface you can also define these material properties with Brooks and Corey's analytic permeability and retention formulas ([Ref. 3](#)) or by interpolation from experimental data.

Richards' equation accounts for changes in the fluid volume fraction with time, and also for changes in the storage related to variations in the pressure head according to Bear ([Ref. 4](#)). With the storage terms, the transport of diluted species in porous media equations in the Subsurface Flow Module also account explicitly for the time changes in liquid and air content.

## Model Definition

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The water moves from a ring on the ground into the subsurface. The 0.25-m radius ring ponds the water to a depth of 0.01 m but is open to the ground surface. Permeable soils exist to a depth of 1.3 m. The soil in the uppermost layer is slightly less permeable than the bottom one. The lower layer sits above relatively impermeable soil, so only a very small amount of leakage exits from the base. The flow is symmetric about the line  $r = 0$ . No flow crosses the surface outside the ring, and there is no flow across the line  $r = 1.25$  m. The initial distribution of pressure heads is known.

The water in the ring contains a dissolved solute at a constant concentration,  $c_0$ . The solute enters the ground with the water and moves through the subsurface by advection and dispersion. Additionally, the solute adsorbs or attaches to solid surfaces, which reduces the aqueous concentrations and also slows solute movement relative to the water. Microbial degradation also reduces both the liquid-phase and solid-phase concentrations. The sorption and the biodegradation are linearly proportional to aqueous concentrations. The fluid in the ring is the only chemical source, and the solute is free to leave the soil column with the fluid flux. Initially the soil is free of the solute. You track its transport for five days.

### FLUID FLOW

Richards' equation governs the saturated-unsaturated flow of water in the soil. The soil air is open to the atmosphere, so you can assume that pressure changes in air do not affect the flow and use Richards' equation here for single-phase flow. Given by [Ref. 4](#), Richards' equation in pressure head reads

$$(C + SeS)\frac{\partial H_p}{\partial t} + \nabla \cdot (-K\nabla(H_p + D)) = 0$$

where  $C$  denotes specific moisture capacity ( $\text{m}^{-1}$ );  $Se$  is the effective saturation of the soil (dimensionless);  $S$  is a storage coefficient ( $\text{m}^{-1}$ );  $H_p$  is the pressure head (m), which is proportional to the dependent variable,  $p$  (Pa);  $t$  is time;  $K$  equals the hydraulic conductivity (m/s);  $D$  is the direction (typically, the  $z$  direction) that represents vertical elevation (m).

To be able to combine boundary conditions and sources with the Darcy's Law formulation, COMSOL Multiphysics converts Richards' equation to SI units and solves for the pressure (SI unit: Pa). Hydraulic head,  $H$ , pressure head,  $H_p$ , and elevation  $D$  are related to pressure  $p$  as

$$H_p = \frac{p}{\rho g}; \quad H = H_p + D$$

Also, the permeability  $\kappa$  (SI unit:  $1/m^2$ ) and hydraulic conductivity  $K$  (SI unit: m/s) are related to the viscosity  $\mu$  (SI unit: Pa s) and density  $\rho$  (SI unit:  $kg/m^3$ ) of the fluid, and the acceleration of gravity  $g$  (SI unit:  $m/s^2$ ) by

$$\frac{\kappa}{\mu} = \frac{K}{\rho g}$$

In this problem,  $S = (\theta_s - \theta_r)/(1 - \rho_r g)$  where  $\theta_s$  and  $\theta_r$  denote the volume fraction of fluid at saturation and after drainage, respectively. For more details see [The Richards' Equation Interface](#) in the *Subsurface Flow Module User's Guide*.

With variably saturated flow, fluid moves through but may or may not completely fill the pores in the soil, and  $\theta$  denotes the volume fraction of fluid within the soil. The coefficients  $C$ ,  $Se$ , and  $K$  vary with the pressure head,  $H_p$ , and with  $\theta$ , making Richards' equation nonlinear. The specific moisture capacity,  $C$ , relates variations in soil moisture to pressure head as  $C = \partial\theta/\partial H_p$ . In the governing equation,  $C$  defines storage changes produced by varying fluid content because  $C\partial H_p/\partial t = \partial\theta/\partial t$ . Because  $C$ , the first term in the time coefficient, goes to zero at saturation, time change in storage relates to compression of the aquifer and water under saturated conditions. The saturated storage comes about with the effective saturation, as represented by the second term in the time-coefficient.

Furthermore,  $K$  is a function that defines how readily the porous media transmits fluid. The relative permeability of the soil,  $k_r$ , increases with fluid content giving  $K = K_s k_r$ , where  $K_s$  (m/d) is the constant hydraulic conductivity at saturation.

This example uses predefined interfaces for van Genuchten formulas ([Ref. 2](#)) to represent how the four retention and permeability properties— $\theta$ ,  $C$ ,  $Se$ , and  $k_r = K/K_s$ —vary with the solution  $H_p$ . The van Genuchten expressions read as follows:

$$\theta = \begin{cases} \theta_r + \text{Se}(\theta_s - \theta_r) & H_p < 0 \\ \theta_s & H_p \geq 0 \end{cases}$$

$$\text{Se} = \begin{cases} \frac{1}{[1 + |\alpha H_p|^n]^m} & H_p < 0 \\ 1 & H_p \geq 0 \end{cases}$$

$$C = \begin{cases} \frac{\alpha m}{1-m} (\theta_s - \theta_r) \text{Se}^{\frac{1}{m}} \left(1 - \text{Se}^{\frac{1}{m}}\right)^m & H_p < 0 \\ 0 & H_p \geq 0 \end{cases}$$

$$k_r = \begin{cases} \text{Se}^l \left[1 - \left(1 - \text{Se}^{\frac{1}{m}}\right)^{m-2}\right] & H_p < 0 \\ 1 & H_p \geq 0 \end{cases}$$

Here  $\alpha$ ,  $n$ ,  $m$ , and  $l$  are constants that specify the type of medium, with  $m = 1 - 1/n$ . In the equations, the system reaches saturation when fluid pressure is atmospheric (that is,  $H_p = 0$ ). When the soil fully saturates, the four parameters reach constant values.

#### *Boundary Conditions and Initial Conditions*

The problem statement records all the boundary conditions you need for this model. The level of water in the ring is known at 0.01 m, giving a Dirichlet constraint on pressure head. Approximate the small leak from the base,  $N_0$ , as  $0.01K_s$ . With no flow crossing the surface outside of the pressure ring or the vertical walls, the following expressions summarize the boundary conditions:

$$\begin{aligned} H_p &= H_{p0} & \partial\Omega \text{ Ring} \\ \mathbf{n} \cdot \mathbf{u} &= 0 & \partial\Omega \text{ Surface} \\ \mathbf{n} \cdot \mathbf{u} &= 0 & \partial\Omega \text{ Sides} \\ \mathbf{n} \cdot \mathbf{u} &= 0 & \partial\Omega \text{ Symmetry} \\ \mathbf{n} \cdot \rho \mathbf{u} &= N_0 & \partial\Omega \text{ Base} \end{aligned}$$

In these expressions,  $\mathbf{n}$  is the unit vector normal to the bounding surface, and  $\mathbf{u}$  is Darcy's velocity field. The initial conditions specify the pressure head in the modeling domain as

$$H_p = \begin{cases} -(z + 1.2 \text{ m}) - 0.2(z + 0.4 \text{ m}) & \text{Upper layer} \\ -(z + 1.2 \text{ m}) & \text{Lower layer} \end{cases}$$

### TRANSPORT OF DILUTED SPECIES IN POROUS MEDIA

Groundwater flow and solute transport are linked by fluid velocities. With the form of the transport equation that follows, the fluid velocities need to come from Darcy's law:

$$\mathbf{u} = K_s k_r \nabla(H_p + D)$$

In this expression,  $\mathbf{u}$  is Darcy's velocity field (SI unit: m/s).

The equation that governs advection, dispersion, sorption, and decay of solutes in groundwater is

$$\frac{\partial}{\partial t}(\theta c) + \frac{\partial}{\partial t}(\rho_b c_P) + \mathbf{u} \cdot \nabla c + \nabla \cdot [-\theta D_L \nabla c] = \Sigma R_L + \Sigma R_P + S_c$$

It describes time rate of change in two terms:  $c$  denotes dissolved concentration (mol/m<sup>3</sup>), and  $c_P$  is the mass of adsorbed contaminant per dry unit weight of solid (mg/kg). Further,  $\theta$  denotes the volume fluid fraction (porosity), and  $\rho_b$  is the bulk density (kg/m<sup>3</sup>). Because  $\rho_b$  amounts to the dried solid mass per bulk volume of the solids and pores together, the term  $\rho_b c_P$  gives solute mass attached to the soil as a concentration. In the equation,  $D_L$  is the hydrodynamic dispersion tensor (m<sup>2</sup>/d);  $R_L$  represents reactions in water (mol/(m<sup>3</sup>·d)); and  $R_P$  equals reactions involving solutes attached to soil particles (mol/(m<sup>3</sup>·d)). Finally,  $S_c$  is solute added per unit volume of soil per unit time (mol/(m<sup>3</sup>·d)).

It is far more convenient to solve the above equation only for dissolved concentration. This requires expanding the left-hand side to

$$\frac{\partial}{\partial t}(\theta c) + \frac{\partial}{\partial t}(\rho_b c_P) = \theta \frac{\partial c}{\partial t} + c \frac{\partial \theta}{\partial t} + \rho_b \frac{\partial c_P}{\partial c} \frac{\partial c}{\partial t}$$

and inserting a few definitions.

In this problem, solute mass per solid mass,  $c_P$ , relates to dissolved concentration,  $c$ , through a linear isotherm or partition coefficient  $k_P$  (m<sup>3</sup>/kg) where  $c_P = k_P c$ . Because the relationship is linear, the derivative is  $k_P = \partial c_P / \partial c$ . Making those substitutions gives the form of the solute transport problem you solve:

$$(\theta + \rho_b k_p) \frac{\partial c}{\partial t} + c \frac{\partial \theta}{\partial t} + \mathbf{u} \cdot \nabla c + \nabla \cdot (-\theta D_L \nabla c) = \theta \phi_L c + \rho_b k_p \phi_P c + S_c$$

In the equation,  $\phi_L$  and  $\phi_P$  denote the decay rates ( $d^{-1}$ ) for the dissolved and adsorbed solute concentrations, respectively.

Select the **Time change in pressure head** option for **Fluid fraction time change** in the **Saturation** settings for the Partially Saturated Porous Media feature to employ results from the flow equation in the solute-transport model:

$$c \frac{\partial \theta}{\partial t} = c C \frac{\partial H_p}{\partial t}$$

Note that COMSOL Multiphysics solves for pressure,  $p$ , and converts to  $H_p$  based on the fluid weight.

The hydrodynamic dispersion tensor,  $D_L$ , describes mechanical spreading from groundwater movement in addition to chemical diffusion:

$$\begin{aligned} \theta D_{Lii} &= \alpha_1 \frac{u_i^2}{|\mathbf{u}|} + \alpha_2 \frac{u_j^2}{|\mathbf{u}|} + \theta \frac{D_m}{\tau_L} \\ \theta D_{Lij} &= \theta D_{Lij} = (\alpha_1 - \alpha_2) \frac{u_i u_j}{|\mathbf{u}|} \end{aligned}$$

where  $D_{Lii}$  are the diagonal entries in the dispersion tensor;  $D_{Lij}$  are the cross terms;  $\alpha$  is the dispersivity (m); the subscripts “1” and “2” denote longitudinal and transverse dispersivities, respectively;  $D_m$  denotes the coefficient of molecular diffusion ( $m^2/d$ ); and  $\tau_L$  is a tortuosity factor that reduces impacts of molecular diffusion for porous media relative to free water. Here  $\tau_L = \theta^{-7/3} \theta_s^2$ .

#### *Boundary Conditions and Initial Conditions*

The boundary and initial conditions in the sorbing-solute problem are straightforward. The solute enters only with the water from the ring at a concentration  $c_0$ . The solute is free to leave, but there is only minimal leakage from the lower boundary and no flow from the sides. Transport is symmetric about the line  $r = 0$ . The boundary conditions in this problem are:

$$\begin{aligned}
c &= c_0 & \partial\Omega \text{ Ring} \\
\mathbf{n} \cdot (-\theta D_L \nabla c) &= 0 & \partial\Omega \text{ Surface} \\
\mathbf{n} \cdot (-\theta D_L \nabla c) &= 0 & \partial\Omega \text{ Sides} \\
\mathbf{n} \cdot (-\theta D_L \nabla c + \mathbf{u}c) &= 0 & \partial\Omega \text{ Symmetry} \\
\mathbf{n} \cdot (-\theta D_L \nabla c) &= 0 & \partial\Omega \text{ Base}
\end{aligned}$$

where  $\mathbf{n}$  is the unit vector normal to the boundary. Because the soil is pristine at the start of the experiment, the initial condition is one of zero concentration.

### MODEL DATA

The following table provides data for the fluid-flow model:

VARIABLE	UNIT	DESCRIPTION	UPPER LAYER	LOWER LAYER
$K_s$	m/d	Saturated hydraulic conductivity	0.298	0.454
$\theta_s$		Porosity/void fraction	0.399	0.339
$\theta_r$		Residual saturation	0.001	0.001
$\alpha$	$m^{-1}$	alpha parameter	1.74	1.39
$n$		n parameter	1.38	1.60
$m$		m parameter	$1-1/n$	$1-1/n$
$l$		Pore connectivity parameter	n/a	
$H_{p0}$	m	Pressure head in ring	0.01	
$H_{p,init}$	m	Initial pressure head	$-(z+1.2)$ $-0.2(z+0.4)$	$-(z+1.2)$

The inputs needed for the solute-transport model are:

VARIABLE	UNITS	DESCRIPTION	VALUE
$\rho_b$	$kg/m^3$	Bulk density	1400
$k_p$	$m^3/kg$	Partition coefficient	0.0001
$D_m$	$m^2/d$	Coefficient of molecular diffusion	0.00374
$\alpha_r$	m	Longitudinal dispersivity	0.005
$\alpha_z$	m	Transverse dispersivity	0.001
$\phi_L$	$d^{-1}$	Decay rate in liquid	0.05
$\phi_P$	$d^{-1}$	Decay rate on solid	0.01
$c_0$	$mol/m^3$	Solute concentration in ring	1.0



## Results and Discussion

Figure 2 and Figure 3 give the solution to the fluid-flow problem at 0.3 days and 1 day, respectively. The images show effective saturation (surface plot), pressure head (contours), and velocities (arrows). The figures illustrate the soil wetting with time. As the arrows indicate, the velocities just below the ring are high relative to the remainder of the soil column.

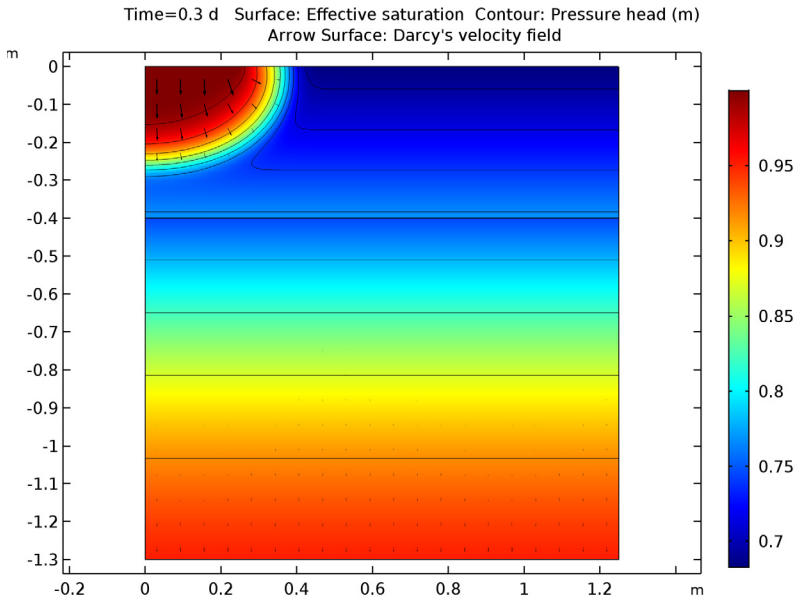
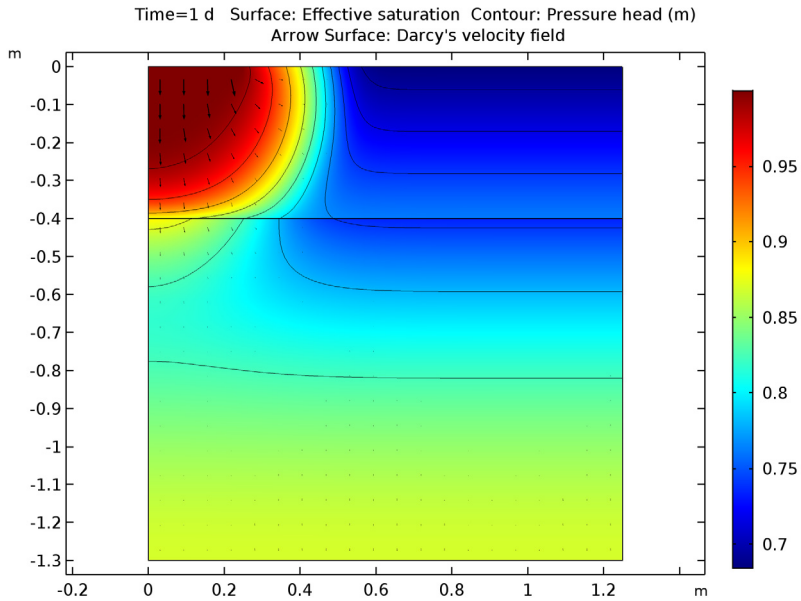
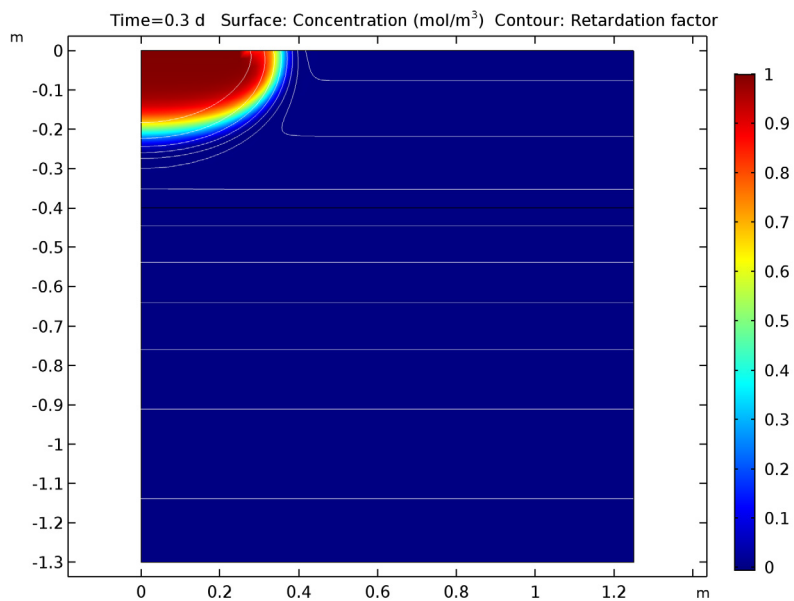


Figure 2: Estimates of effective saturation (surface plot), pressure head (contours), and velocity (arrows) in variably saturated soil after 0.3 days.



*Figure 3: Estimates of effective saturation (surface plot), pressure head (contours), and velocity (arrows) in variably saturated soil after 1 day.*

Figure 4 and Figure 5 give the concentrations for 0.3 days and 1 day, respectively, along with the retardation factor. They illustrate how the solute concentrations (surface plot) enter and move through the soil. Because the retardation factor depends on soil moisture, its value varies with the solution.



*Figure 4: Solution for dissolved concentrations (surface plot) and retardation factor (contours) at 0.3 days.*

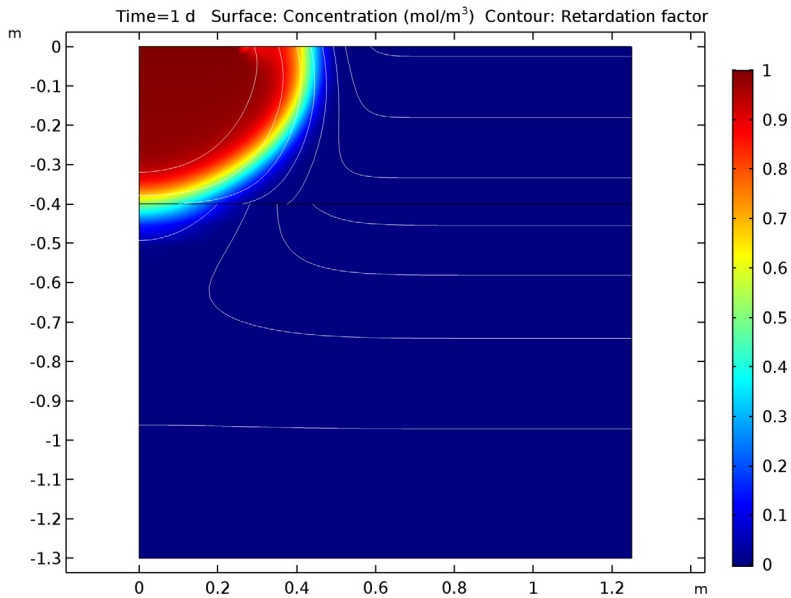


Figure 5: Solution for dissolved concentrations (surface plot) and retardation factor (contours) at 1 day.

Figure 6 shows an image of the retardation factor at the end of the simulation time interval. For variably saturated solute transport, the retardation factor changes with time. As shown in this image, the process of sorption has the greatest potential to slow the contaminant where the soils are relatively dry. The retardation coefficient here ranges from roughly 1.35 to 1.55, and the solute moves at approximately the velocity of the groundwater.

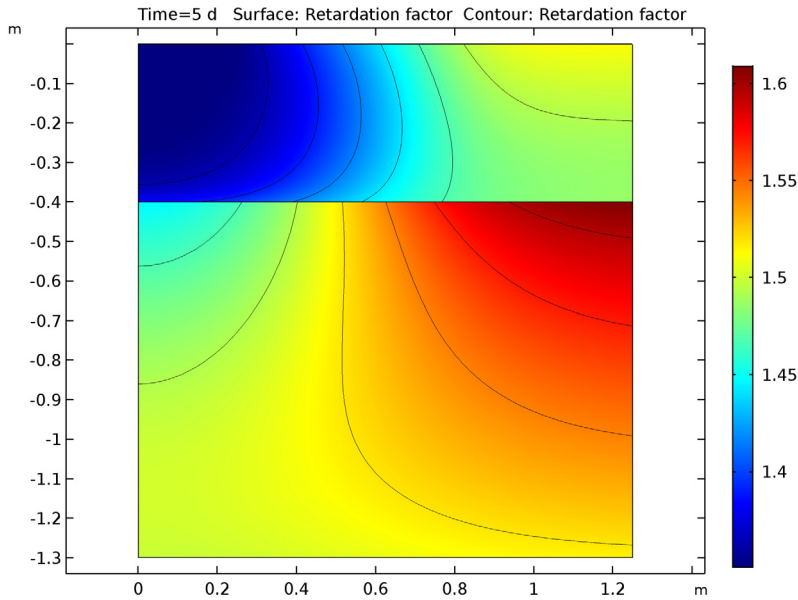


Figure 6: Snapshot of the retardation factor (surface and contours) at 5 days.

### *Notes About the COMSOL Implementation*

This model makes use of the **Infinite Element Domain** feature. It performs a coordinate scaling to the selected domain such that boundary conditions on the outside of the infinite element layer are effectively applied at a very large distance. Therefore unwanted effects of artificial boundary conditions on the region of interest are suppressed. This allows to model details in an area which is actually very large or infinite.

### *References*

1. J. Simunek, T. Vogel, and M.Th. van Genuchten, “The SWMS\_2D code for simulating water flow and solute transport in two-dimensional variably saturated media,” ver. 1.1., *Research Report No. 132*, U.S. Salinity Laboratory, USDA, 1994.
2. M.Th. van Genuchten, “A closed-form equation for predicting the hydraulic of conductivity of unsaturated soils,” *Soil Sci. Soc. Am. J.*, vol. 44, pp. 892–898, 1980.

3. R.H. Brooks and A.T. Corey, “Properties of porous media affecting fluid flow,” *J. Irrig. Drainage Div.*, ASCE Proc. 72(IR2), pp. 61–88, 1966.

4. J. Bear, *Hydraulics of Groundwater*, McGraw-Hill, 1978.

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**Application Library path:** Subsurface\_Flow\_Module/Solute\_Transport/  
sorbing\_solute

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### *Modeling Instructions*

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From the **File** menu, choose **New**.

#### **NEW**

In the **New** window, click **Model Wizard**.

#### **MODEL WIZARD**

- 1 In the **Model Wizard** window, click **2D Axisymmetric**.
- 2 In the **Select Physics** tree, select **Fluid Flow>Porous Media and Subsurface Flow>Richards’ Equation (dl)**.
- 3 Click **Add**.
- 4 In the **Select Physics** tree, select **Chemical Species Transport>Transport of Diluted Species in Porous Media (tds)**.
- 5 Click **Add**.
- 6 Click **Study**.
- 7 In the **Select Study** tree, select **Preset Studies for Selected Physics Interfaces>Time Dependent**.
- 8 Click **Done**.

#### **GLOBAL DEFINITIONS**

Load the parameters from file.

##### *Parameters*

- 1 On the **Home** toolbar, click **Parameters**.
- 2 In the **Settings** window for **Parameters**, locate the **Parameters** section.
- 3 Click **Load from File**.

- 4 Browse to the model's Application Libraries folder and double-click the file `sorbing_solute_parameters.txt`.

## GEOMETRY I

The modeling domain is made up of the two permeable soil layers, each of which is represented by a rectangular domain in 2D axisymmetry.

### *Rectangle 1 (r1)*

- 1 On the **Geometry** toolbar, click **Primitives** and choose **Rectangle**.
- 2 In the **Settings** window for **Rectangle**, locate the **Size and Shape** section.
- 3 In the **Width** text field, type 1.5.
- 4 In the **Height** text field, type 0.9.
- 5 Locate the **Position** section. In the **z** text field, type -1.3.
- 6 Click to expand the **Layers** section. Select the **Layers to the right** check box.
- 7 Clear the **Layers on bottom** check box.
- 8 In the table, enter the following settings:

Layer name	Thickness (m)
Layer 1	0.25

This additional layer to the right will later be used to define an **Infinite Element Domain**. Read more about it in the [Notes About the COMSOL Implementation](#) section. Proceed with the second soil layer.

### *Rectangle 2 (r2)*

- 1 On the **Geometry** toolbar, click **Primitives** and choose **Rectangle**.
- 2 In the **Settings** window for **Rectangle**, locate the **Size and Shape** section.
- 3 In the **Width** text field, type 1.5.
- 4 In the **Height** text field, type 0.4.
- 5 Locate the **Position** section. In the **z** text field, type -0.4.
- 6 Locate the **Layers** section. Select the **Layers to the right** check box.
- 7 Clear the **Layers on bottom** check box.

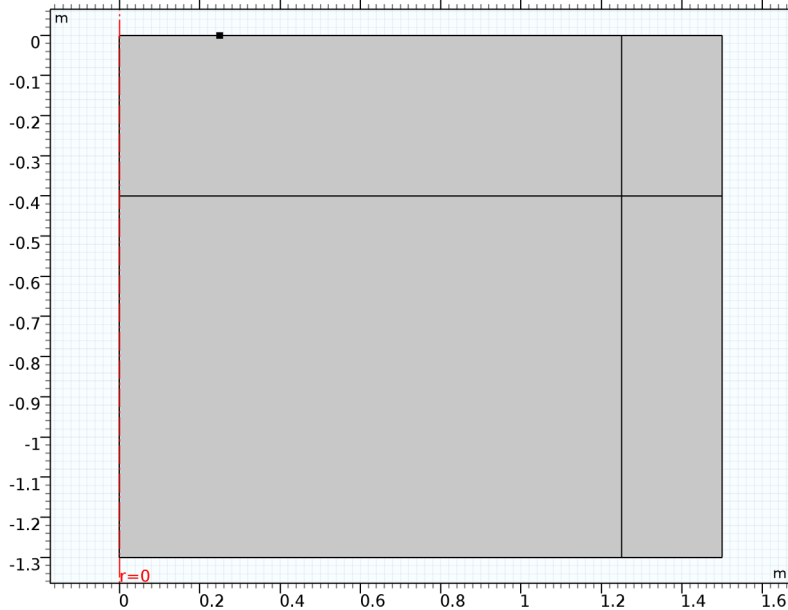
8 In the table, enter the following settings:

Layer name	Thickness (m)
Layer 1	0.25

To finish the model geometry, add a point on the top boundary marking the pond's outer rim.

*Point 1 (pt1)*

- 1 On the **Geometry** toolbar, click **Primitives** and choose **Point**.
- 2 In the **Settings** window for **Point**, locate the **Point** section.
- 3 In the **r** text field, type 0.25.
- 4 Click **Build All Objects**.
- 5 Click the **Zoom Extents** button on the **Graphics** toolbar.



Now, define the **Infinite Element Domain**.

## DEFINITIONS

*Infinite Element Domain 1 (ie1)*

- 1 On the **Definitions** toolbar, click **Infinite Element Domain**.
- 2 Select Domains 3 and 4 only.



- 3 In the **Settings** window for **Infinite Element Domain**, locate the **Geometry** section.
- 4 From the **Type** list, choose **Cylindrical**.

### **RICHARDS' EQUATION (DL)**

Begin by specifying the properties for the bottom soil layer in the default Richards' Equation Model node, then duplicate this node and modify the domain selection and properties to match the top layer.

#### *Richards' Equation Model 1*

- 1 In the **Model Builder** window, under **Component 1 (comp1)>Richards' Equation (dl)** click **Richards' Equation Model 1**.
- 2 In the **Settings** window for **Richards' Equation Model**, locate the **Fluid Properties** section.
- 3 From the  $\rho$  list, choose **User defined**. In the associated text field, type rho.
- 4 Locate the **Matrix Properties** section. From the **Permeability model** list, choose **Hydraulic conductivity**.
- 5 In the  $K_s$  text field, type Ks\_1.
- 6 In the  $\theta_s$  text field, type thetas\_1.
- 7 In the  $\theta_r$  text field, type thetar\_1.
- 8 Locate the **Storage Model** section. From the **Storage** list, choose **User defined**. In the  $S$  text field, type Ss\_1.
- 9 Locate the **Retention Model** section. In the  $\alpha$  text field, type alpha\_1.
- 10 In the  $n$  text field, type n\_1.

#### *Richards' Equation Model 2*

- 1 Right-click **Component 1 (comp1)>Richards' Equation (dl)>Richards' Equation Model 1** and choose **Duplicate**.
- 2 Select Domains 2 and 4 only.
- 3 In the **Settings** window for **Richards' Equation Model**, locate the **Matrix Properties** section.
- 4 In the  $K_s$  text field, type Ks\_2.
- 5 In the  $\theta_s$  text field, type thetas\_2.
- 6 In the  $\theta_r$  text field, type thetar\_2.
- 7 Locate the **Storage Model** section. In the  $S$  text field, type Ss\_2.
- 8 Locate the **Retention Model** section. In the  $\alpha$  text field, type alpha\_2.
- 9 In the  $n$  text field, type n\_2.

#### *Initial Values 1*

- 1 In the **Model Builder** window, under **Component 1 (comp1)>Richards' Equation (dl)** click **Initial Values 1**.
- 2 In the **Settings** window for **Initial Values**, locate the **Initial Values** section.
- 3 Click the **Pressure head** button.
- 4 In the  $H_p$  text field, type  $-(z+1.2)$ .

#### *Initial Values 2*

- 1 Right-click **Component 1 (comp1)>Richards' Equation (dl)>Initial Values 1** and choose **Duplicate**.
- 2 Select Domains 2 and 4 only.
- 3 In the **Settings** window for **Initial Values**, locate the **Initial Values** section.
- 4 In the  $H_p$  text field, type  $-(z+1.2) - 0.2*(z+0.4)$ .

#### *Pressure Head 1*

- 1 On the **Physics** toolbar, click **Boundaries** and choose **Pressure Head**.
- 2 Select Boundary 5 only.
- 3 In the **Settings** window for **Pressure Head**, locate the **Pressure Head** section.
- 4 In the  $H_{p0}$  text field, type  $H_{p0}$ .

#### *Pervious Layer 1*

- 1 On the **Physics** toolbar, click **Boundaries** and choose **Mass Flux**.
- 2 On the **Physics** toolbar, click **Boundaries** and choose **Pervious Layer**.
- 3 Select Boundaries 2 and 8 only.
- 4 In the **Settings** window for **Pervious Layer**, locate the **Pervious Layer** section.
- 5 In the  $H_b$  text field, type  $-2$ .
- 6 In the  $R_b$  text field, type  $1/5[d]$ .

#### *Gravity 1*

- 1 In the **Model Builder** window, under **Component 1 (comp1)>Richards' Equation (dl)** click **Gravity 1**.
- 2 In the **Settings** window for **Gravity**, locate the **Gravity** section.
- 3 From the **Specify** list, choose **Elevation**.

### **TRANSPORT OF DILUTED SPECIES IN POROUS MEDIA (TDS)**

Now, set up the transport equation for an unsaturated porous medium, accounting for dispersion and adsorption

On the **Physics** toolbar, click **Richards' Equation (dl)** and choose **Transport of Diluted Species in Porous Media (tds)**.

- 1 In the **Model Builder** window, under **Component 1 (comp1)** click **Transport of Diluted Species in Porous Media (tds)**.
- 2 In the **Settings** window for **Transport of Diluted Species in Porous Media**, locate the **Transport Mechanisms** section.
- 3 Find the **Porous media transport** subsection. Select the **Dispersion** check box.

*Partially Saturated Porous Media I*

- 1 On the **Physics** toolbar, click **Domains** and choose **Partially Saturated Porous Media**.
- 2 In the **Settings** window for **Partially Saturated Porous Media**, locate the **Domain Selection** section.
- 3 From the **Selection** list, choose **All domains**.
- 4 Locate the **Convection** section. From the  **$\mathbf{u}$**  list, choose **Darcy's velocity field (dl)**.
- 5 Locate the **Matrix Properties** section. From the  $\epsilon_p$  list, choose **User defined**. In the associated text field, type `d1.theta`.

This corresponds to the saturated liquid volume fraction, defined in the **Richards' Equation** interface.

- 6 From the  $\rho$  list, choose **User defined**. In the associated text field, type `rhob`.
- 7 Locate the **Saturation** section. From the list, choose **Liquid volume fraction**.
- 8 In the  $\theta$  text field, type `d1.theta`.

This corresponds to the liquid volume fraction.

- 9 From the **Fluid fraction time change** list, choose **Time change in pressure head**.
- 10 From the  $dH_p/dt$  list, choose **Time change in pressure head (dl)**.
- 11 In the  $C_m$  text field, type `d1.Cm`.
- 12 Locate the **Diffusion** section. In the  $D_{L,c}$  text field, type `D1`.
- 13 Locate the **Dispersion** section. From the **Dispersion tensor** list, choose **Dispervivity**.
- 14 From the **Dispervivity model** list, choose **Transverse isotropic**.
- 15 In the  $\alpha$  table, enter the following settings:

---

`alphar`

---

`alphaz`

---

Add adsorption as a subnode to the **Partially Saturated Porous Media** node.

### *Adsorption 1*

- 1 On the **Physics** toolbar, click **Attributes** and choose **Adsorption**.
- 2 In the **Settings** window for **Adsorption**, locate the **Adsorption** section.
- 3 From the **Species c** list, choose **User defined**.
- 4 In the  $k_{p,c}$  text field, type  $kp$ .

### *Reactions 1*

- 1 On the **Physics** toolbar, click **Domains** and choose **Reactions**.
- 2 In the **Settings** window for **Reactions**, locate the **Domain Selection** section.
- 3 From the **Selection** list, choose **All domains**.
- 4 Locate the **Reaction Rates** section. In the  $R_c$  text field, type  $(\text{phip} * kp * \text{rhub} - \text{phil} * \text{dl.theta}) * c$ .

### *Outflow 1*

- 1 On the **Physics** toolbar, click **Boundaries** and choose **Outflow**.
- 2 Select Boundaries 2, 8, 12, and 13 only.

### *Concentration 1*

- 1 On the **Physics** toolbar, click **Boundaries** and choose **Concentration**.
- 2 Select Boundary 5 only.
- 3 In the **Settings** window for **Concentration**, locate the **Concentration** section.
- 4 Select the **Species c** check box.
- 5 In the  $c_{0,c}$  text field, type  $c0$ .

## **MESH 1**

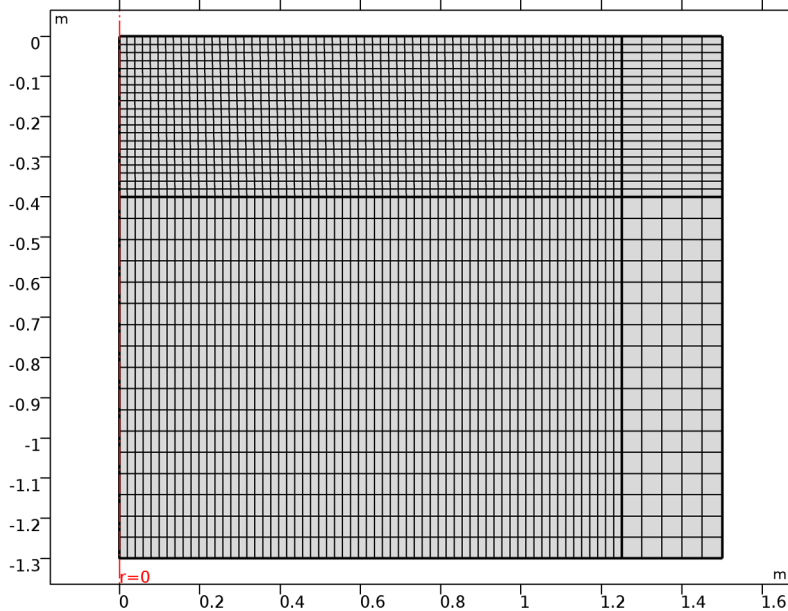
Using a mapped mesh is a good idea for this geometry. It uses less mesh elements while keeping the accuracy compared to using a triangular mesh with the same mesh size.

- 1 In the **Model Builder** window, under **Component 1 (comp1)** click **Mesh 1**.
- 2 In the **Settings** window for **Mesh**, locate the **Mesh Settings** section.
- 3 From the **Element size** list, choose **Finer**.

### *Size 1*

- 1 Right-click **Component 1 (comp1)>Mesh 1** and choose **Mapped**.
- 2 In the **Model Builder** window, under **Component 1 (comp1)>Mesh 1** right-click **Mapped 1** and choose **Size**.
- 3 In the **Settings** window for **Size**, locate the **Geometric Entity Selection** section.

- 4 From the **Geometric entity level** list, choose **Domain**.
- 5 Select Domain 2 only.
- 6 Locate the **Element Size** section. Click the **Custom** button.
- 7 Locate the **Element Size Parameters** section. Select the **Maximum element size** check box.
- 8 In the associated text field, type 0.02.
- 9 In the **Model Builder** window, click **Mesh 1**.
- 10 In the **Settings** window for **Mesh**, click **Build All**.



## STUDY 1

### Step 1: Time Dependent

- 1 In the **Settings** window for **Time Dependent**, locate the **Study Settings** section.
- 2 From the **Time unit** list, choose **d**.
- 3 In the **Times** text field, type range(0,0.1,0.9) range(1,1,5).
- 4 On the **Home** toolbar, click **Compute**.

## RESULTS

### *Data Sets*

Flownet, pressure and concentration plots are created per default. Pressure and concentration are also visualized on a revolved 3D geometry. Visualizing the results on the infinite element domains doesn't add value to the plots. Focus on the region close to the source and therefore hide the infinite element domains from the plots with the following steps.

### *Study 1/Solution 1 (sol1)*

In the **Model Builder** window, expand the **Results>Data Sets** node, then click **Study 1/Solution 1 (sol1)**.

### *Selection*

- 1 On the **Results** toolbar, click **Selection**.
- 2 In the **Settings** window for **Selection**, locate the **Geometric Entity Selection** section.
- 3 From the **Geometric entity level** list, choose **Domain**.
- 4 Select Domains 1 and 2 only.

### *Pressure (dl)*

The second default plot group contains a surface plot of the pressure distribution. Modify it to show the effective saturation, pressure head, and velocity field at different times.

- 1 In the **Model Builder** window, under **Results** click **Pressure (dl)**.
- 2 In the **Settings** window for **2D Plot Group**, type **Effective saturation** in the **Label** text field.
- 3 Locate the **Data** section. From the **Time (d)** list, choose **0.3**.

### *Surface*

- 1 In the **Model Builder** window, expand the **Results>Effective saturation** node, then click **Surface**.
- 2 In the **Settings** window for **Surface**, click **Replace Expression** in the upper-right corner of the **Expression** section. From the menu, choose **Model>Component 1>Richards' Equation>dl.Se - Effective saturation**.
- 3 Click to expand the **Title** section. From the **Title type** list, choose **Custom**.
- 4 Find the **Type and data** subsection. Clear the **Unit** check box.

### *Contour 1*

- 1 In the **Model Builder** window, under **Results** right-click **Effective saturation** and choose **Contour**.

- 2 In the **Settings** window for **Contour**, click **Replace Expression** in the upper-right corner of the **Expression** section. From the menu, choose **Model>Component 1>Richards' Equation>dl.Hp - Pressure head**.
- 3 Locate the **Levels** section. In the **Total levels** text field, type 10.
- 4 Locate the **Coloring and Style** section. From the **Coloring** list, choose **Uniform**.
- 5 From the **Color** list, choose **Black**.
- 6 Clear the **Color legend** check box.
- 7 Click to expand the **Quality** section. From the **Resolution** list, choose **Finer**.

#### *Arrow Surface 1*

- 1 Right-click **Effective saturation** and choose **Arrow Surface**.
- 2 In the **Settings** window for **Arrow Surface**, locate the **Coloring and Style** section.
- 3 From the **Color** list, choose **Black**.
- 4 Locate the **Arrow Positioning** section. Find the **R grid points** subsection. In the **Points** text field, type 20.
- 5 Find the **Z grid points** subsection. In the **Points** text field, type 20.
- 6 On the **Effective saturation** toolbar, click **Plot**.

Compare the plot in the Graphics window with that in [Figure 2](#).

#### *Effective saturation*

- 1 In the **Model Builder** window, under **Results** click **Effective saturation**.
- 2 In the **Settings** window for **2D Plot Group**, locate the **Data** section.
- 3 From the **Time (d)** list, choose **1**.
- 4 On the **Effective saturation** toolbar, click **Plot**.

Compare with the plot in [Figure 3](#).

The third default plot shows the solute concentration. Follow the steps below to reproduce the plots shown in [Figure 4](#) and [Figure 5](#).

#### *Concentration (tds)*

- 1 In the **Model Builder** window, under **Results** click **Concentration (tds)**.
- 2 In the **Settings** window for **2D Plot Group**, locate the **Data** section.
- 3 From the **Time (d)** list, choose **0.3**.

#### *Contour 1*

- 1 Right-click **Results>Concentration (tds)** and choose **Contour**.

- 2 In the **Settings** window for **Contour**, click **Replace Expression** in the upper-right corner of the **Expression** section. From the menu, choose **Model>Component 1>Transport of Diluted Species in Porous Media>tds.RF\_c - Retardation factor**.
- 3 Click to expand the **Title** section. From the **Title type** list, choose **Custom**.
- 4 Find the **Type and data** subsection. Clear the **Unit** check box.
- 5 Locate the **Levels** section. In the **Total levels** text field, type 10.
- 6 Locate the **Coloring and Style** section. From the **Coloring** list, choose **Uniform**.
- 7 From the **Color** list, choose **White**.
- 8 Clear the **Color legend** check box.
- 9 Locate the **Quality** section. From the **Resolution** list, choose **Fine**.
- 10 On the **Concentration (tds)** toolbar, click **Plot**.

Compare the result with that in [Figure 4](#).

#### *Concentration (tds)*

- 1 In the **Model Builder** window, under **Results** click **Concentration (tds)**.
- 2 In the **Settings** window for **2D Plot Group**, locate the **Data** section.
- 3 From the **Time (d)** list, choose **1**.
- 4 On the **Concentration (tds)** toolbar, click **Plot**.

Compare with [Figure 5](#).

Finally, plot the retardation factor after 5 days ([Figure 6](#)).

#### *Concentration (tds) 2*

- 1 Right-click **Results>Concentration (tds)** and choose **Duplicate**.
- 2 In the **Settings** window for **2D Plot Group**, type Retardation factor in the **Label** text field.
- 3 Locate the **Data** section. From the **Time (d)** list, choose **5**.

#### *Surface 1*

- 1 In the **Model Builder** window, expand the **Results>Retardation factor** node, then click **Surface 1**.
- 2 In the **Settings** window for **Surface**, click **Replace Expression** in the upper-right corner of the **Expression** section. From the menu, choose **Model>Component 1>Transport of Diluted Species in Porous Media>tds.RF\_c - Retardation factor**.
- 3 Locate the **Title** section. From the **Title type** list, choose **Custom**.
- 4 Find the **Type and data** subsection. Clear the **Unit** check box.



*Contour 1*

- 1** In the **Model Builder** window, under **Results>Retardation factor** click **Contour 1**.
- 2** In the **Settings** window for **Contour**, locate the **Coloring and Style** section.
- 3** From the **Color** list, choose **Black**.
- 4** On the **Retardation factor** toolbar, click **Plot**.

